

Ocean Mixing

as an Emerging **Essential Ocean Variable** (EOV)

Arnaud Le Boyer (SIO), **Nicole Couto** (SIO),
Laura Cimolli (SIO), Bruce Howe (UH)



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Ocean Mixing

Drivers, Mechanisms and Impacts



“the three-dimensional turbulent interleaving and blending of oceanic waters with different properties.”

**Ocean Mixing as an
Essential Ocean
Variable**

Le Boyer, Couto, et. al
(in prep.)

Ocean Mixing as an emerging EOVS



Relevance: The variable is effective in addressing the overall GOOS Themes – Climate, Real-Time Services, and Ocean Health.



Feasibility: Observing or deriving the variable on a *global scale* is technically feasible using proven, scientifically understood methods.

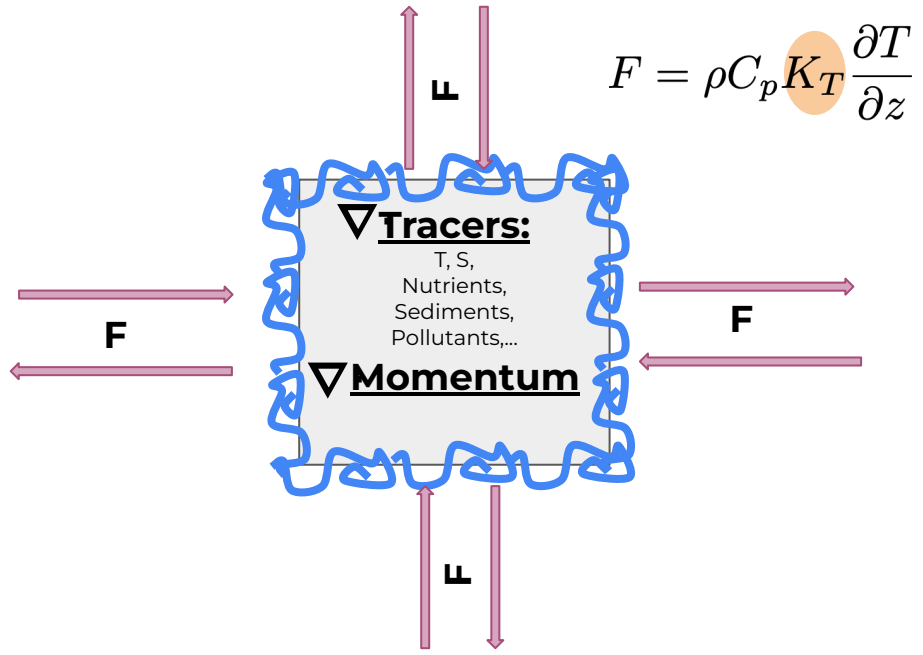


Cost effectiveness: Generating and *archiving data* on the variable is affordable, mainly relying on *coordinated observing systems* using proven technology, taking advantage where possible of historical datasets.

What is ocean mixing?

Ocean mixing refers to turbulent processes occurring at small spatial (1 m) and temporal (seconds to few minutes) scales that lead to the diabatic transformation of water masses.

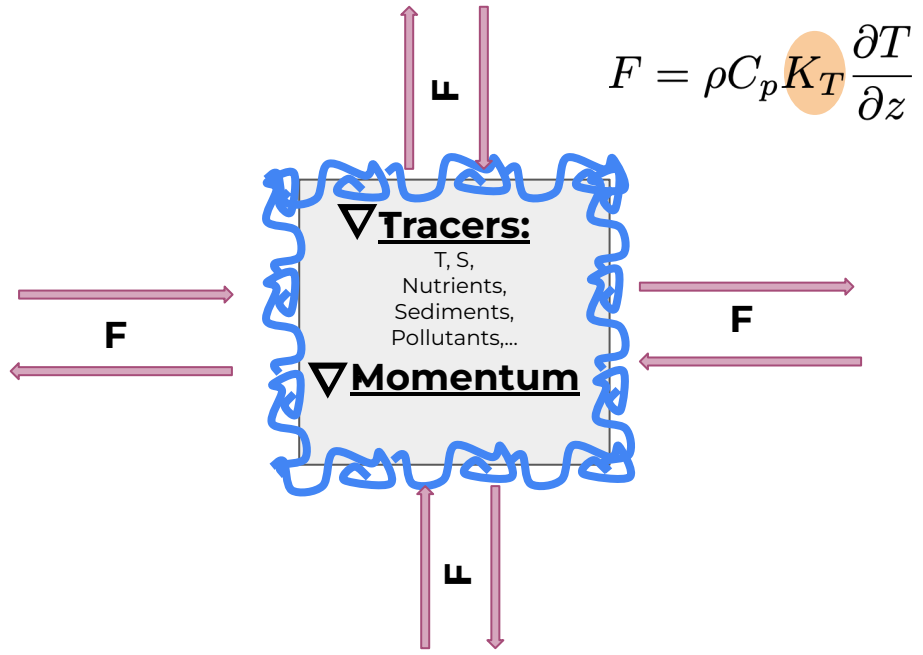
Mathematically, we describe this with **turbulent fluxes** - the rate at which a tracer moves across a boundary.



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Diffusivity describes how readily temperature, salinity, dissolved nutrients, etc can move across isopycnals in the ocean.

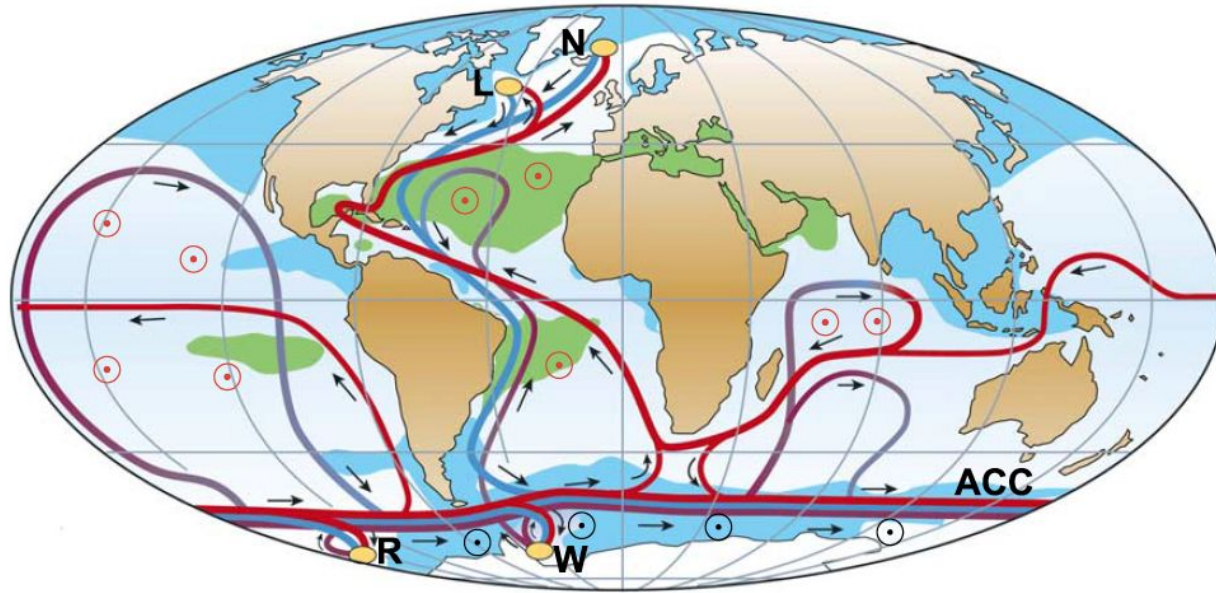
Mixing rates describe how quickly energy is dissipated.

$$K_{\rho} = \frac{\Gamma \langle \epsilon \rangle}{\langle N^2 \rangle}$$

$$K_T = \frac{\langle \chi \rangle}{6 \langle \frac{\partial T}{\partial z}^2 \rangle}$$

Scientific Relevance

- The upwelling branch of the thermohaline circulation relies on turbulent mixing.



— Surface flow
— Deep flow
— Bottom flow
● Deep Water Formation

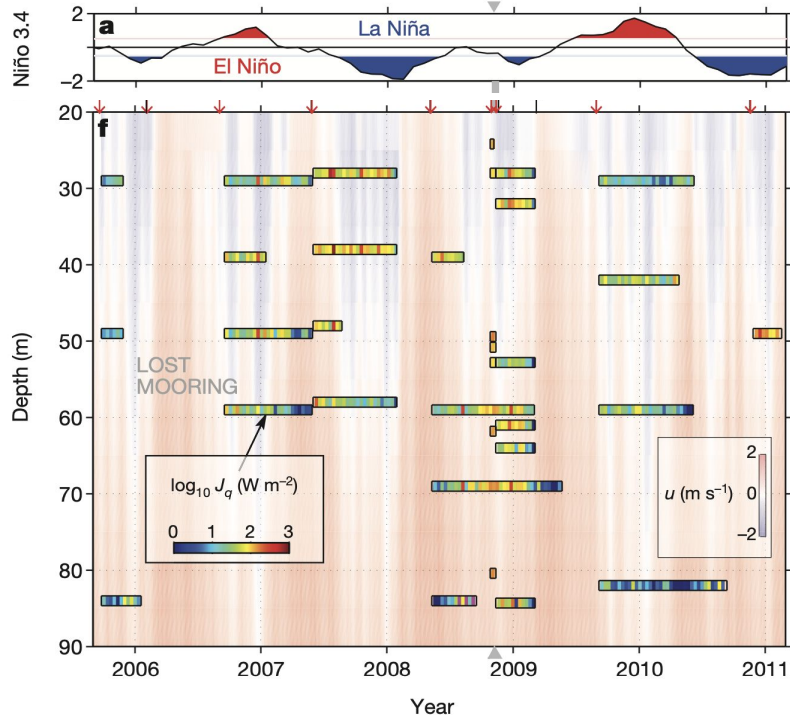
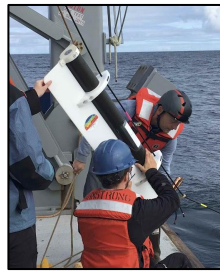
⊙ Wind-driven upwelling
⊙ Mixing-driven upwelling
■ Salinity > 36 ‰
■ Salinity < 34 ‰

L Labrador Sea
N Nordic Seas
W Weddell Sea
R Ross Sea

Scientific Relevance

- Many climate models fail to accurately simulate the seasonal cycle of the sea surface temperature in the equatorial cold tongue.

χ Pods on moorings at the TAO array

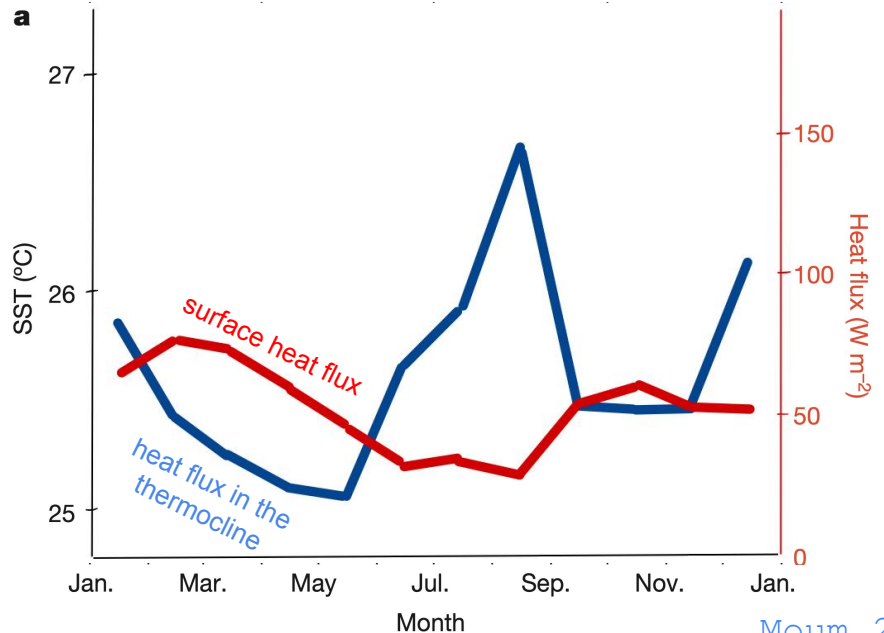
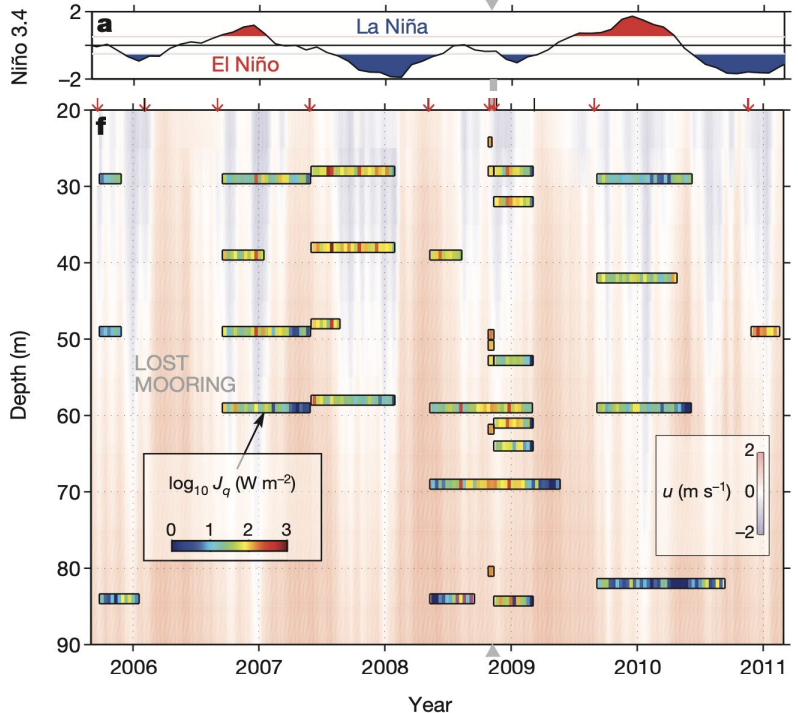
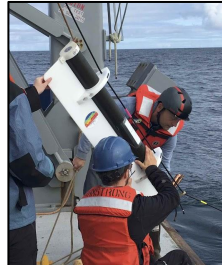


- Moum et al used a multi-year timeseries of heat flux computed from χ Pods to show that SST is related to mixing in the thermocline.

Scientific Relevance

- Jan-Jun: surface heat flux > flux across thermocline
- Jul-Dec: flux across thermocline > surface heat flux

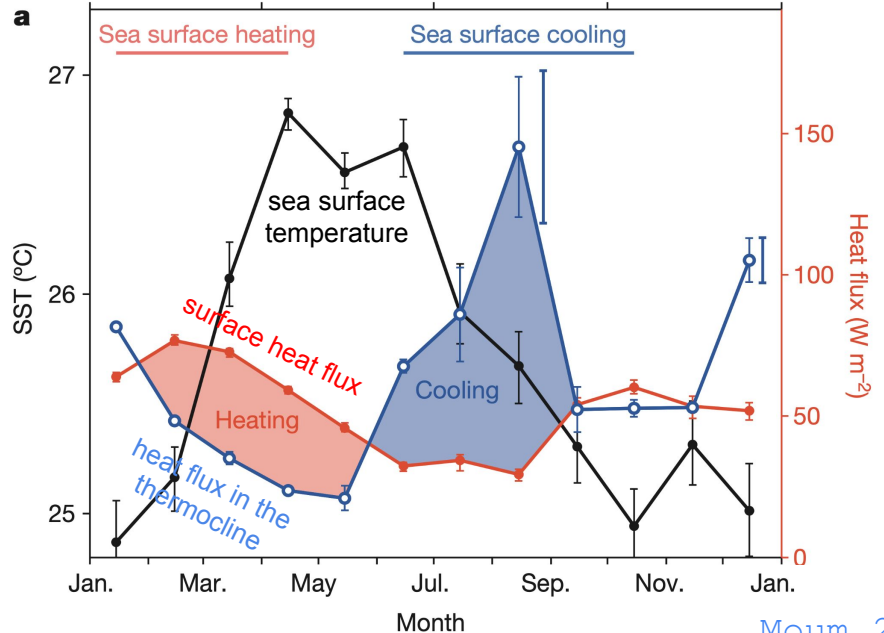
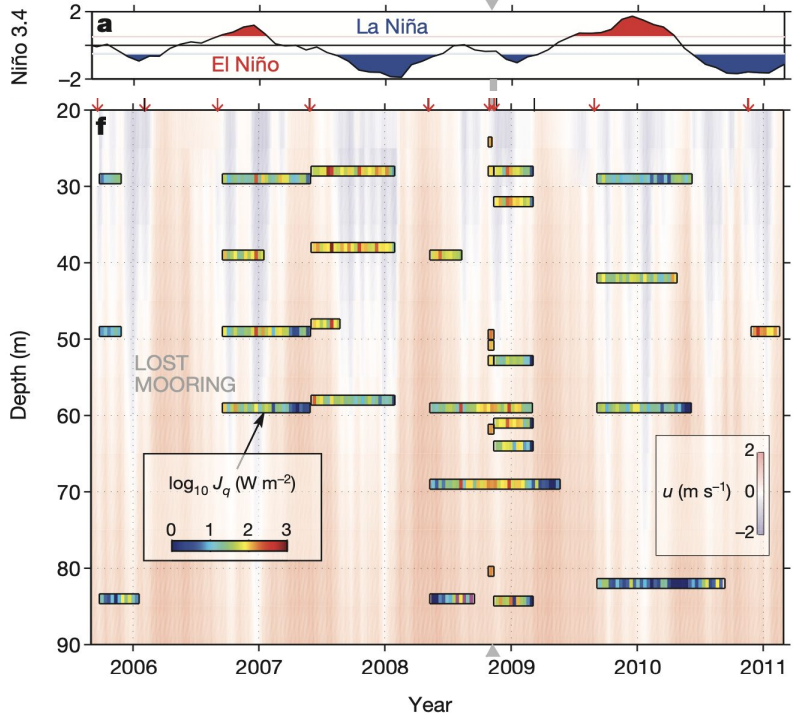
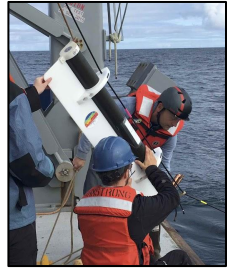
xpods on moorings at the TAO array



Scientific Relevance

- Equatorial SST warms when atmospheric heating exceeds cooling from mixing below, and cools when mixing exceeds atmospheric heating.

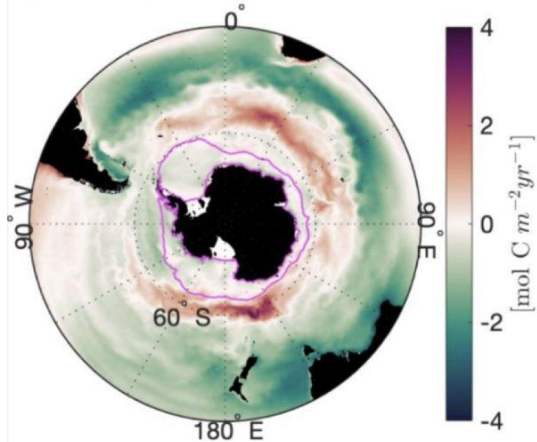
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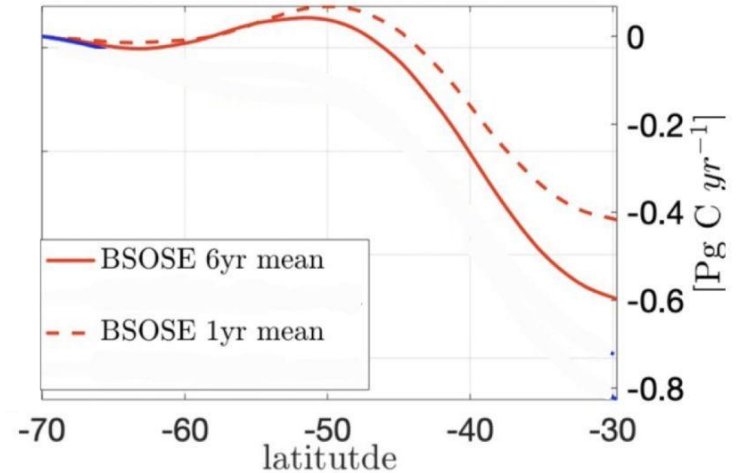
Scientific Relevance :

- Turbulent mixing has a significant impact on air-sea fluxes

(a) CO_2 flux, BSOSE

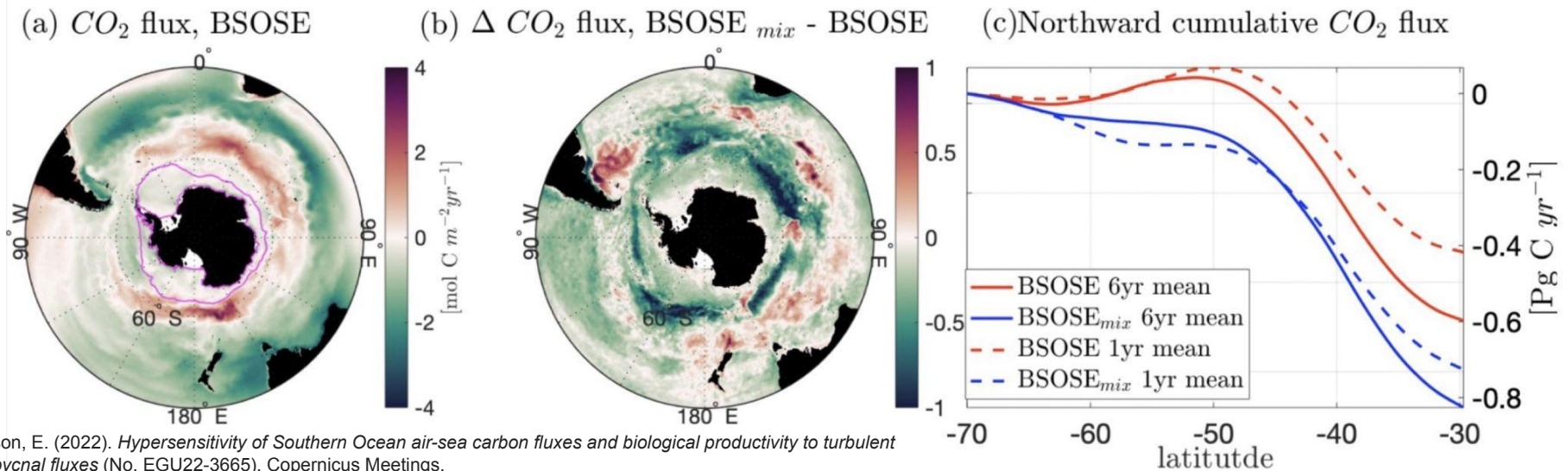


(c) Northward cumulative CO_2 flux

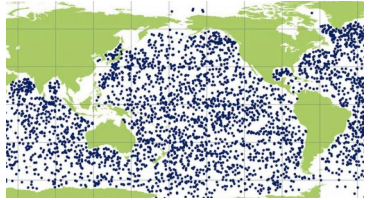
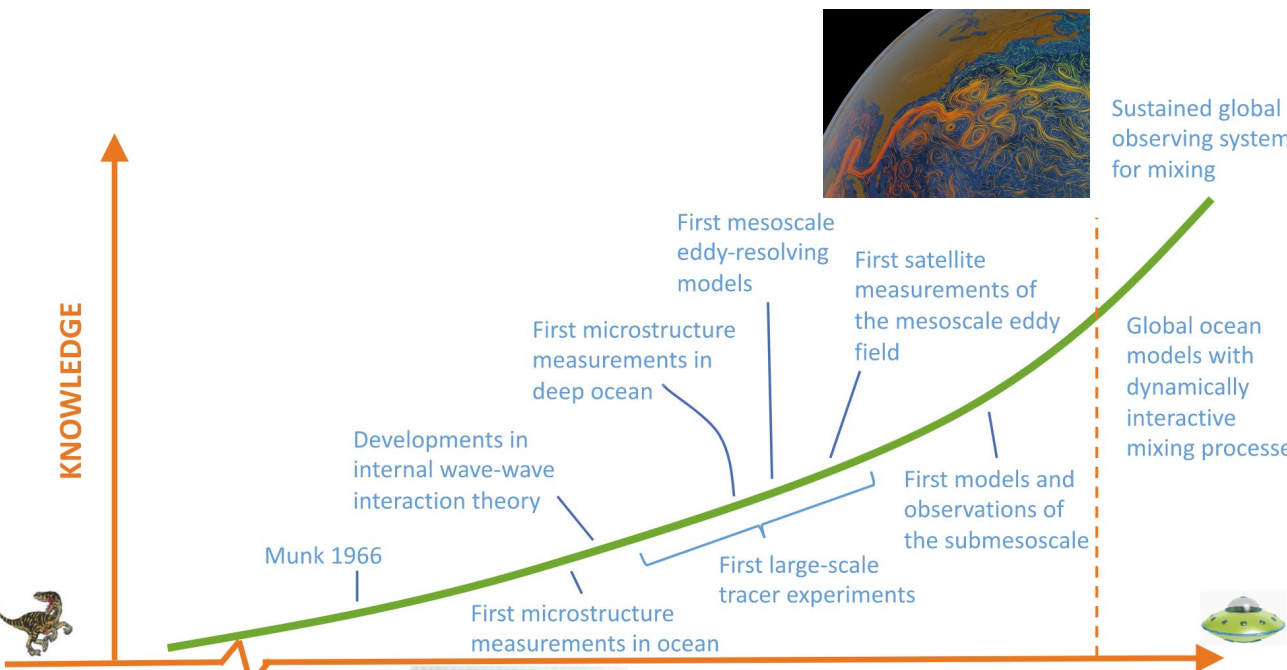


Scientific Relevance

- Turbulent mixing has a significant impact on air-sea fluxes
- Altering the mixing in an eddy-permitting ocean model alters the distribution of inorganic carbon, alkalinity, temperature, and salinity
- These changes alter the biological productivity, which affects the total carbon flux



Feasibility / Technological maturity



- Turbulence sensors have become less expensive and more user-friendly
- Increasing in-situ and remote observations, combined with increasing computation power has led to better models and parameterizations
- Now, the community is focusing on moving toward sustained global observations

The dawn of ocean science

1960s

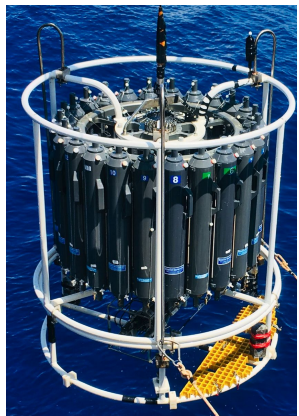
Today



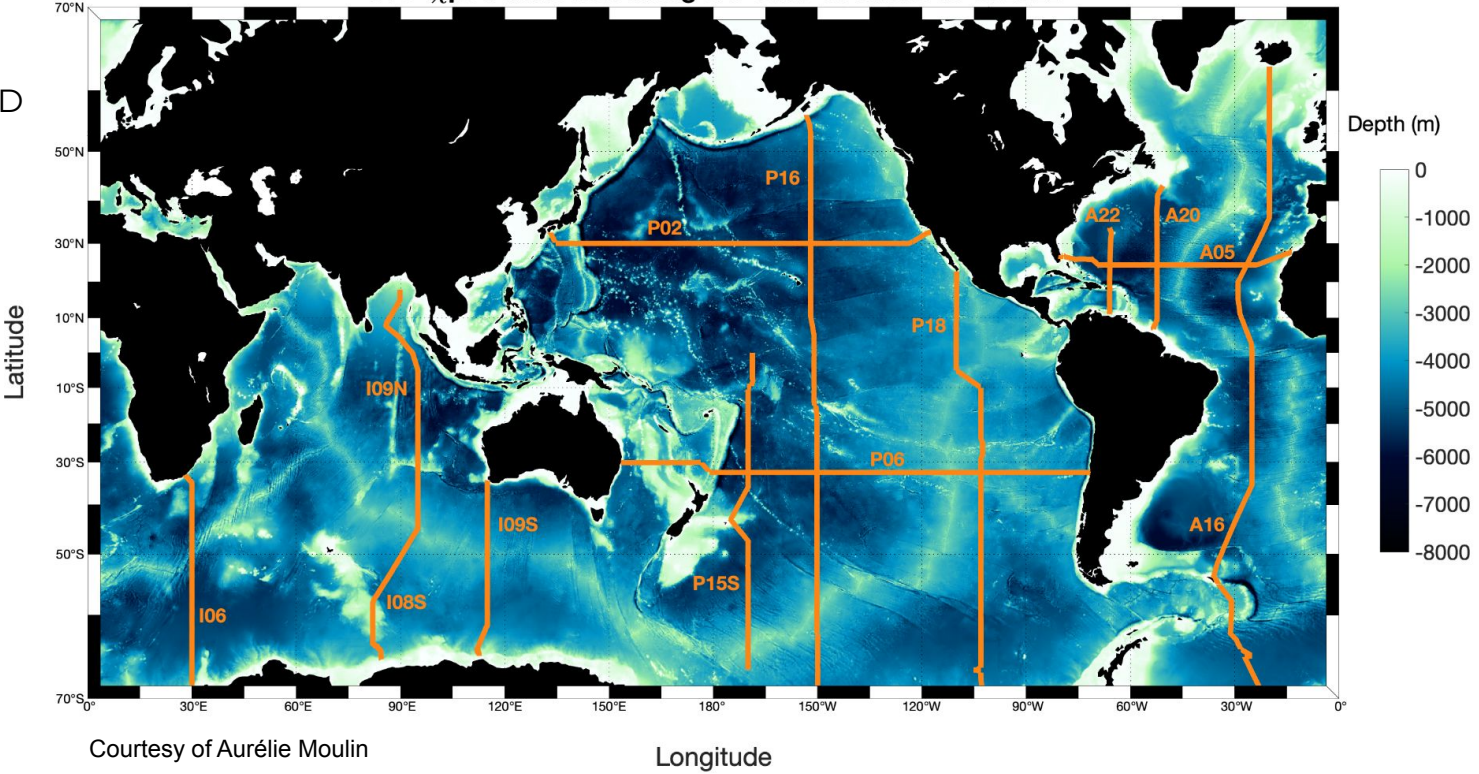
Coordinated efforts

GO-SHIP program:

χ profiles can be measured with χ pods attached to CTD rosette.



CTD- χ pod transects along GO-SHIP sections as of 2022



Courtesy of Aurélie Moulin

Longitude

Coordinated efforts

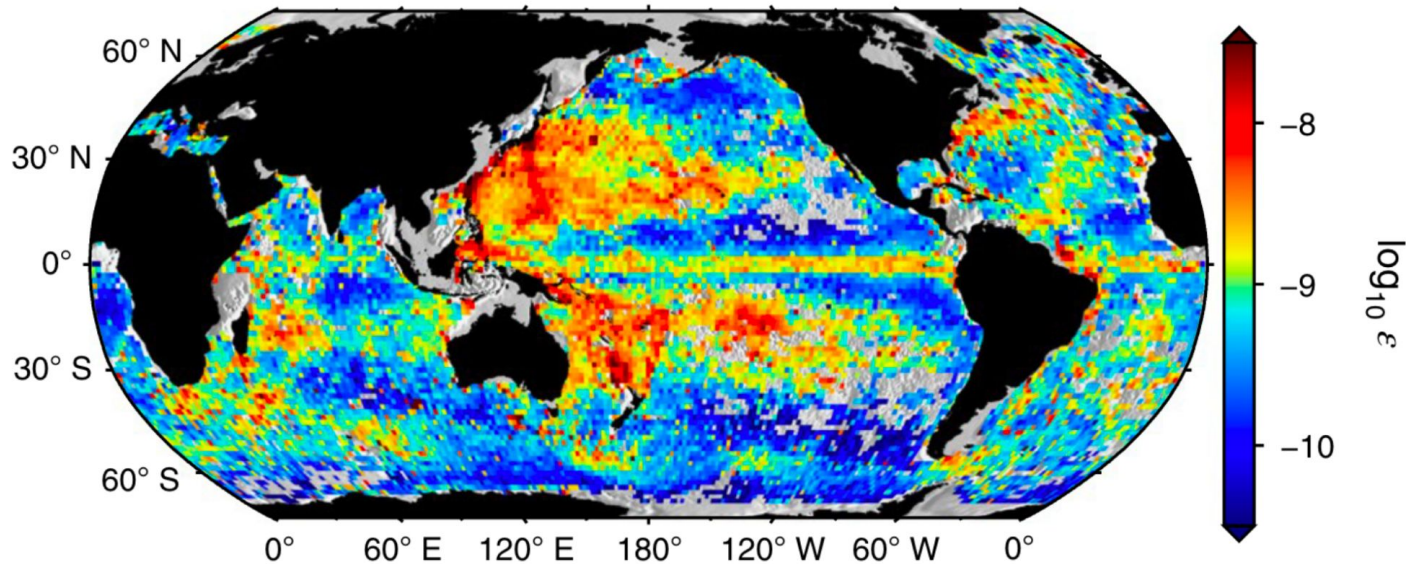
ARGO network:

can make global estimates of dissipation from density profiles and strain parameterizations.



b

Average dissipation rate 250–500 m (W kg^{-1})

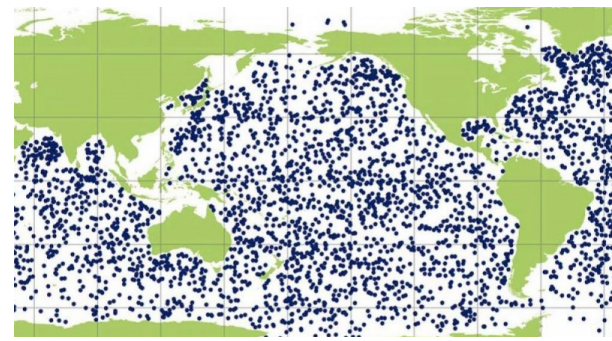
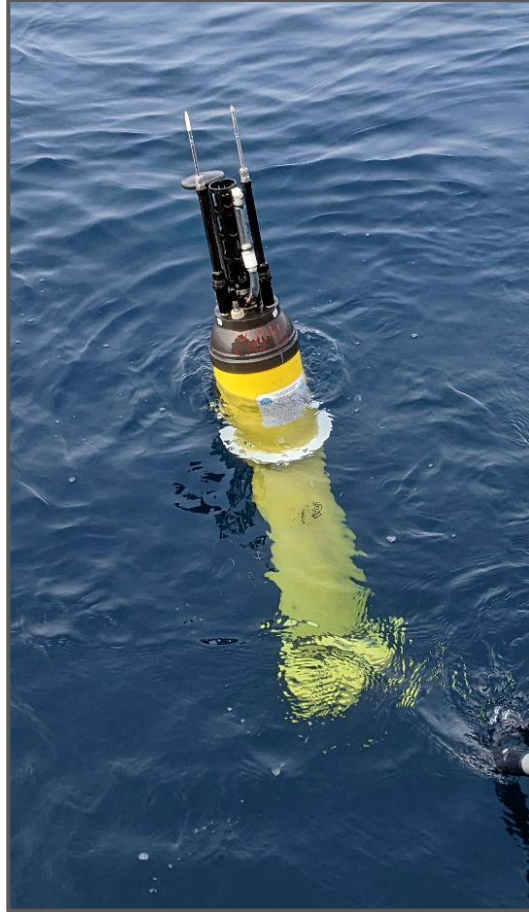


Coordinated efforts







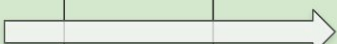
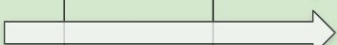





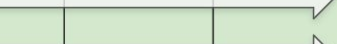



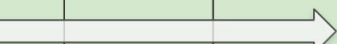
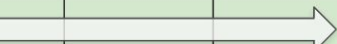

ARGO network:

can make global estimates of dissipation from density profiles and strain parameterizations.

ARGO-MIX: can directly measure dissipation rates with new on-board processing capabilities



Based on Readiness Levels, Ocean Mixing should be considered an emerging EOV

	Level 1 Idea	Level 2 Documentation	Level 3 Proof of concept	Level 4 Trial	Level 5 Verification	Level 6 Operational	Level 7 Fitness for purpose	Level 8 Mission qualified	Level 9 Sustained
Requirements			  			ADVs Tethered profilers On moorings (e.g. γ -pods) ADCPs Gliders AUVs (propelled)	        		
Coordination			 	GO-SHIP Argo-MIX			 		
Data management					Open data repositories ATOMIX SCOR		 		

Ocean Bottom Pressure

An advantage of OBP sensors is their ability to resolve time/space scales through regional arrays and the variability of large-scale circulation on time scales of *days*, *weeks*, and *months* (especially at latitudes $>40^\circ$ where surface/deep flows are coupled)

For periods <1 day, OBP is sensitive even to weak

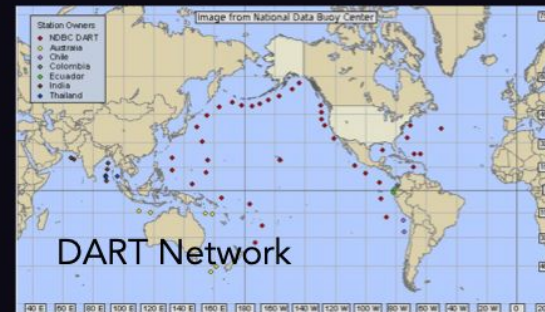
- gravitational tides
- internal tides
- internal waves
- tsunamis and other infragravity waves
- storm surges

Observing tsunamis is a critical task for GOOS!

Determining accurate tidal constants, including secular changes, is possible due to low geophysical “noise” at the seafloor

- essential to de-aliasing and correcting other measurements (e.g. those from satellite altimetry)
- provides essential information about structure of Earth’s crust and vertical deformation of the seafloor

Requirements Settings*					
Responsible GCOS/GOOS Panel Reporting Mechanism	OOPC GCOS Implementation Plan/Status Reporting to UNFCCC				
Readiness Level ⁵	Mature Level 8. [Several basin-scale and regional telemetering networks fully operational, e.g., DART; OOI RCN; ONC; DONET; ACO.]				
Phenomena to Capture	Sea Level	Circulation	Fronts and Eddies	Tides	Infragravity Waves and Tsunamis
Temporal Scales of the Phenomena (order)	Week to decade	Week to decade	Day to month	Hour to week	Minute to hour
Horizontal Scales of the Phenomena (order)	10 km to 10,000 km	10 km to 10,000 km	1 km to 100 km	10 km to 10,000 km	1 km to 1,000 km
Magnitudes of the Signals of the Phenomena (order)	0.01 dbar to 1.0 dbar	0.01 dbar to 1.0 dbar	0.01 dbar to 1.0 dbar	0.001 dbar to 1.0 dbar	0.001 dbar to 1.0 dbar





Ocean Bottom Pressure

- Ocean bottom pressure (OBP) is the total pressure resulting from the mass of the column of seawater and the overlying atmosphere
 - Fluctuations in the atmospheric sea level pressure (the dynamic topography) and fluctuations in the mass of the fluid column between the unperturbed surface and bottom contribute to the observed OBP variability
 - OBP serves as a proxy of ocean mass variability
 - ↳ a key parameter needed in many geophysical applications (e.g. large-scale ocean circulation - important for climate) or variability in the Earth's gravity field
 - OBP observations can provide estimates of both basin-wide modes of transport between OBP sensors located on opposing continental slopes and local bottom current amplitudes related to local energetic circulation
 - ↳ OBP observations are needed to improve the accuracy of global ocean circulation and are a key parameter needed to calibrate and improve remote sensing missions that measure the mass changes in the ocean like Gravity Recovery and Climate Experiment (GRACE)

Ocean Bottom Pressure

Examples of available networks and datasets

Observation Deployment & Maintenance				
Observing Elements ⁶	DART Stations	DONET, OOI RCN, ONC, ACO	Moorings (OceanSITES, DBCP)	GRACE
Relevant measured parameter	Pressure	Pressure	Pressure	Gravity perturbations
Measurement Type	in-situ	in-situ	in-situ	remote
Sensor(s)/Technique	SBE 54 Tsunami Pressure Sensor	SBE 54 and Paroscientific Nano-resolution Digiquartz Pressure Transducer	Paroscientific Digiquartz Pressure Transducer, e.g., incorporated in PIES instrument./	Inversion of relative positions and speeds of paired satellites.
Phenomena addressed	Infragravity waves Tsunamis Tides Sea level	Infragravity waves Tsunamis Tides Eddies Sea level	Tides Fronts and eddies Circulation Sea Level	Circulation Sea Level
Readiness Level	Mature Level 8	Mature Level 8	Mature Level 8	Mature Level 9
Spatial sampling	Specific locations; 100's of km spacing and up	Specific locations; 1 km to 100's of km spacing	Point samples; fixed locations at tens of km spacing and up	100's of km
Temporal sampling	15 s to 15 min	1 Hz	Better than 1 Hz to several samples per hour	Monthly gridded maps

Scales resolved by current technologies

